A SLOPE STABILITY RELIABILITY MODEL

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Keywords: slope stability, reliability, spatial variability

Abstract

Evaluating the reliability of slopes against sliding failure is complicated by the fact that most slope soils are spatially variable. This means that instead of nice circular failure surfaces, slope failures tend to be more complex, following the weakest path or zones through the material.

The finite element method is well suited to slope stability calculations since it allows the failure surface to seek out the weakest path through the soil. This paper wedds the finite element method with random field simulations to perform a ‘random finite element method’ (RFEM) analysis of a slope. The simulation results are used to validate a simplified analytical reliability model for slope stability which is based on harmonic averaging. The harmonic average is low-strength dominated, so that it captures the weakest path characteristics of slope failure and provides good estimates of the failure probability.

1. INTRODUCTION

The failure prediction of a soil slope has been a long-standing geotechnical problem, and one which has attracted a wide variety of solutions. Traditional approaches to the problem generally involve assuming that the soil slope is homogeneous (spatially constant), or possibly layered, and techniques such as Taylor’s (1937) stability coefficients for frictionless soils, the method of slices, and other more general methods involving arbitrary failure surfaces have been developed over the years. The main drawback to these methods is that they are not able to easily find the critical failure surface in the event that the soil properties are spatially varying.

In the realistic case where the soil properties vary randomly in space, the slope stability problem is best captured via a non-linear finite element model which has the distinct advantage of allowing the failure surface to seek out the path of least resistance. In this paper such a model is employed, which, when combined with a random field simulator, allows the realistic probabilistic evaluation of slope stability. This work is a follow up of a comparison of probabilistic slope stability methods carried out by Griffiths and Fenton (2004) which looked in some detail at the probability of failure of a single slope geometry. Two slope geometries are considered in this paper, one shallower with a 2:1 gradient and the other steeper with a 1:1 gradient. Both slopes are assumed to be composed of undrained clay, with \( \phi_u = 0 \), of height \( H \) with the slope resting on a foundation layer, also of depth \( H \). The finite element mesh for the 2:1 gradient slope is shown in Figure 1. The 1:1 slope is similar.

The soil is represented by a random spatially varying undrained cohesion field, \( c_u(x) \), which is assumed to be lognormally distributed, where \( x \) is the spatial position. The cohesion has mean \( \mu_{c_u} \), standard deviation \( \sigma_{c_u} \) and is assumed to have an exponentially decaying (Markovian) correlation structure,

\[
\rho_{\ln c_u} (\tau) = e^{-\frac{2|\tau|}{\theta_{\ln c_u}}}
\]

where \( \tau \) is the distance between two points in the field. Note that the correlation structure has been assumed isotropic in this study. The use of an anisotropic correlation is straightforward, within the framework developed
here, but is considered a site specific extension. In this paper it is desired to investigate the stochastic behaviour of slope stability for the simpler isotropic case, leaving the effect of anisotropy for future work.

![Mesh used for stability analysis of the 2:1 gradient slope.](image)

Figure 1. Mesh used for stability analysis of the 2:1 gradient slope.

The correlation function has a single parameter, \( \theta_{\ln c_u} \), the correlation length. Because \( c_u \) is assumed to be lognormally distributed, its logarithm, \( \ln c_u \), is normally distributed. In this study, the correlation function is measured relative to the underlying normally distributed field. Thus, \( \rho_{\ln c_u}(\tau) \) gives the correlation coefficient between \( \ln c_u(x) \) and \( \ln c_u(x') \) at two points in the field separated by the distance \( \tau = |x - x'| \). In practice, the parameter \( \theta_{\ln c_u} \) can be estimated from spatially distributed \( c_u \) samples by using the logarithm of the samples rather than the raw data themselves. If the actual correlation between points in the \( c_u \) field is desired, the following transformation can be used (Vanmarcke, 1984),

\[
\rho_{c_u}(\tau) = \frac{\exp[\rho_{\ln c_u}(\tau)\sigma_{\ln c_u}^2] - 1}{\exp[\sigma_{\ln c_u}^2] - 1} \tag{2}
\]

Since \( \theta_{\ln c_u} \) is a length, it can be non-dimensionalized by dividing it by \( H \), a measure of the embankment size. Thus, the results given here can be applied to any size problem, so long as it has the same slope and same depth to height ratio. The standard deviation, \( \sigma_{c_u} \) may also be expressed in terms of the dimensionless coefficient of variation

\[
V = \frac{\sigma_{c_u}}{\mu_{c_u}} \tag{3}
\]

If the mean and variance of the underlying \( \ln c_u \) field are desired, they can be obtained through the transformations

\[
\sigma_{\ln c_u}^2 = \ln (1 + V^2), \quad \mu_{\ln c_u} = \ln(\mu_{c_u}) - \frac{1}{2}\sigma_{\ln c_u}^2 \tag{4}
\]

By using Monte Carlo simulation, where the soil slope is simulated and analyzed by the finite element method repeatedly, estimates of the probability of failure are obtained over a range of soil statistics. The failure probabilities are compared to those obtained using a harmonic average of the cohesion field employed in Taylor’s stability coefficient method and very good agreement is found. The study indicates that the stability of a spatially varying soil slope is well modeled using a harmonic average of the soil properties.
2. THE RANDOM FINITE ELEMENT MODEL

The slope stability analyses use an elastic-perfectly plastic stress-strain law with a Tresca failure criterion. Plastic stress redistribution is accomplished using a viscoplastic algorithm which uses 8-node quadrilateral elements and reduced integration in both the stiffness and stress redistribution parts of the algorithm. The theoretical basis of the method is described more fully in Chapter 6 of the text by Smith and Griffiths (2004), and for a discussion of the method applied to slope stability analysis, the reader is referred to Griffiths and Lane (1999) and Paice and Griffiths (1997).

In brief, the analyses involve the application of gravity loading, and the monitoring of stresses at all the Gauss points. If the Tresca criterion is violated, the program attempts to redistribute those stresses to neighboring elements that still have reserves of strength. This is an iterative process which continues until the Tresca criterion and global equilibrium are satisfied at all points within the mesh under quite strict tolerances.

In this study, “failure” is said to have occurred if, for any given realization, the algorithm is unable to converge within 500 iterations. Following a set of 2000 realizations of the Monte-Carlo process the probability of failure is simply defined as the proportion of these realizations that required 500 or more iterations to converge.

While the choice of 500 as the iteration ceiling is subjective, Griffiths and Fenton (2000) found that the probability of failure computed using this criterion is quite stable even for as few as 200 iterations, which is to say that convergence of an unfailed slope generally occurs well before 200 iterations.

The random finite element model (RFEM) combines the deterministic finite element analysis with a random field simulator, which, in this study, is the Local Average Subdivision (LAS) method developed by Fenton and Vanmarcke (1990). The LAS algorithm produces a field of random element values, each representing a local average of the random field over the element domain, which are then mapped directly to the finite elements. The random elements are local averages of the log-cohesion, \( \ln c_u \) field. The resulting realizations of the log-cohesion field have correlation structure and variance correctly accounting for local averaging over each element. Much discussion of the relative merits of various methods of representing random fields in finite element analysis has been carried out in recent years (see, for example, Li and Der Kiureghian, 1993). While the spatial averaging discretization of the random field used in this study is just one approach to the problem, it is appealing in the sense that it reflects the simplest idea of the finite element representation of a continuum as well as the way that soil samples are typically taken and tested in practice, i.e. as local averages. Regarding the discretization of random fields for use in finite element analysis, Matthies et al. (1997) makes the comment that “One way of making sure that the stochastic field has the required structure is to assume that it is a local averaging process.”, refering to the conversion of a nondifferentiable to a differentiable (smooth) stochastic process. Matthies further goes on to say that the advantage of the local average representation of a random field is that it yields accurate results even for rather coarse meshes.

Figure 2 illustrates two possible realizations arising from the RFEM for the 2:1 slope – similar results were observed for the 1:1 slope. In this figure, dark regions correspond to stronger soil. Notice how convoluted the failure region is, particularly at the smaller correlation length. It can be seen that the slope failure involves the plastic deformation of a region around a failure ‘surface’ which undulates along the weakest path. Clearly failure is more complex than just a rigid ‘circular’ region sliding along a clearly defined interface, as is typically assumed.
3. PARAMETRIC STUDIES

To keep the study non-dimensional, the mean soil strength is expressed in the form of a mean dimensionless shear strength,

\[ \mu_{N_s} = \frac{\mu_{cu}}{\gamma H} \]  

(5)

where \( \gamma \) is the unit weight of the soil, assumed in this study to be deterministic. In the 2:1 slope case where the cohesion field is assumed to be everywhere the same and equal to \( \mu_{cu} \), a value of \( \mu_{N_s} = 0.173 \) corresponds to a factor-of-safety, \( F = 1.0 \), which is to say that the slope is on the verge of failure. For the 1:1 slope, \( \mu_{N_s} = 0.184 \) corresponds to a factor-of-safety of \( F = 1.0 \). Both of these values were determined by finding the deterministic value of \( c_u \) needed to just achieve failure in the finite element model, bearing in mind that the failure surface cannot descend below the base of the model. These values are almost identical to what would be identified using Taylor’s charts (see also Baker, 2003), although as will be seen later, small variations in the choice of the critical values of \( \mu_{N_s} \) can result in significant changes in the estimated probability of slope failure, particularly for larger factors of safety.

This study considers the following values of the input statistics. For the 2:1 slope, \( \mu_{N_s} \) is varied over the following values:

\[ \mu_{N_s} = 0.15, 0.17, 0.20, 0.25, 0.30 \]

and over

\[ \mu_{N_s} = 0.15, 0.18, 0.20, 0.25, 0.30 \]

for the 1:1 slope. For both slopes, the following ranges in the (normalized) correlation length, \( \theta_{\ln cu}/H \), and coefficient of variation, \( V \), were investigated;

\[ \theta_{\ln cu}/H = 0.10, 0.20, 0.50, 1.00, 2.00, 5.00 \]

\[ V = 0.10, 0.20, 0.50, 1.00, 2.00, 5.00 \]

For each set of the above parameters, 2000 realizations of the soil field were simulated and analyzed, from which the probability of slope failure was estimated. Details of the probability estimates for the 2:1 slope are presented in Griffiths and Fenton (2000). This paper concentrates on the development of a failure probability model, using a harmonic average of the soil, and compares the simulated probability estimates to those predicted by the harmonic average model.
4. SEMI-THEORETICAL MODEL

In Taylor’s stability coefficient approach to slope stability, the coefficient

\[ N_s = \frac{c_u}{\gamma H} \]  \hspace{1cm} (6)

assumes that the soil is completely uniform, having cohesion equal to \( c_u \) everywhere. This coefficient may then be compared to the critical coefficient obtained from Taylor’s charts to determine if slope failure will occur or not. For the slope geometry studied here, slope failure will occur if

\[ N_s < N_{crit} \]

where \( N_{crit} = 0.173 \) for the 2:1 slope and \( N_{crit} = 0.184 \) for the 1:1 slope.

In the case where \( c_u \) is randomly varying in space, two issues present themselves. First of all Taylor’s method cannot be used on a non-uniform soil and secondly Eq. (6) now includes a random quantity on the right-hand-side (namely, \( c_u = c_u(x) \)) so that \( N_s \) becomes random. The first issue can be solved by finding some representative value of \( c_u \), which will be referred to here as \( \bar{c}_u \), such that the stability coefficient method still holds. That is, \( \bar{c}_u \) would be the cohesion of a uniform soil such that it has the same slope stability as the real spatially varying soil.

The question now is, how should this effective soil cohesion value be defined? First of all, each soil realization will have a different value of \( \bar{c}_u \), so that Eq. (6) is still a function of a random quantity, namely,

\[ N_s = \frac{\bar{c}_u}{\gamma H} \]  \hspace{1cm} (7)

If the distribution of \( \bar{c}_u \) is found, the distribution of \( N_s \) can be derived. The failure probability of the slope then becomes equal to the probability that \( N_s \) is less than the Taylor critical value, \( N_{crit} \).

This line of reasoning suggests that \( \bar{c}_u \) should be defined as some sort of average of \( c_u \) over the soil domain where failure is occurring. Three common types of averages present themselves;

1) **Arithmetic average**: the arithmetic average over some domain, \( A \), is defined as,

\[ X_a = \frac{1}{n} \sum_{i=1}^{n} c_{u_i} = \frac{1}{A} \int_A c_u(x) \, dx \]  \hspace{1cm} (8)

for the discrete and continuous cases, where the domain \( A \) is assumed to be divided up into \( n \) samples in the discrete case. The arithmetic average weights all of the values of \( c_u \) equally. In that the failure surface seeks a path through the weakest parts of the soil, this form of averaging is not deemed to be appropriate for this problem.

2) **Geometric average**: the geometric average over some domain, \( A \), is defined as,

\[ X_g = \left( \prod_{i=1}^{n} c_{u_i} \right)^{1/n} = \exp \left\{ \frac{1}{A} \int_A \ln c_u(x) \, dx \right\} \]  \hspace{1cm} (9)

The geometric average is dominated by low values of \( c_u \) and, for a spatially varying cohesion field, will always be less than the arithmetic average. This average potentially reflects the reduced strength as seen along the failure path and has been found by the authors (Fenton and Griffiths, 2002 and 2003) to well represent the bearing capacity and settlement of footings founded on spatially random soils. The geometric average is also a ‘natural’ average of the lognormal distribution, since an arithmetic average of the underlying normally distributed random variable, \( \ln c_u \), leads to the geometric average when converted back to the lognormal distribution. Thus, if \( c_u \) is lognormally distributed, its geometric local average will also be lognormally distributed with the median preserved.

3) **Harmonic average**: the harmonic average over some domain, \( A \), is defined as,

\[ X_h = \left[ \frac{1}{n} \sum_{i=1}^{n} \frac{1}{c_{u_i}} \right]^{-1} = \left[ \frac{1}{A} \int_A \frac{dx}{c_u(x)} \right]^{-1} \]  \hspace{1cm} (10)
This average is even more strongly influenced by small values than is the geometric average. In general, for a spatially varying random field, the harmonic average will be smaller than the geometric average, which in turn is smaller than the arithmetic average. Unfortunately, the mean and variance of the harmonic average, for a spatially correlated random field, are not easily found.

Putting aside for the moment the issue of how to compute the effective undrained cohesion, \( \bar{c}_u \), the size of the averaging domain must also be determined. This should approximately equal the area of the soil which fails during a slope subsidence. Since the value of \( \bar{c}_u \) changes only slowly with changes in the averaging domain, only an approximate area need be determined. The area selected in this study is a parallelogram, as shown in Figure 3, having slope length equal to the length of the slope and horizontal surface length equal to \( H \). For the purposes of computing the average, it is further assumed that this area can be approximated by a rectangle of dimension \( w \times h \) (averages over rectangles are generally easier to compute). Thus, a rectangular \( w \times h \) area is used to represent a roughly circular band (on average) within which the soil is failing in shear.

![Figure 3. Assumed averaging domain (1:1 slope is similar).](image)

In this study, the values of \( w \) and \( h \) are taken to be

\[
w = H / \sin \beta, \quad h = H \sin \beta
\]

such that \( w \times h = H^2 \), where \( \beta \) is the slope angle (26.6° for the 2:1 slope and 45° for the 1:1 slope). It appears, when comparing Figure 2 to Figure 3, that the assumed averaging domain of Figure 3 is smaller than the deformed regions seen in Figure 2. A general prescription for the size of the averaging domain is not yet known, although it should capture the approximate area of the soil involved in resisting the slope deformation. The area assumed in Figure 3 is to be viewed as an initial approximation which, as will be seen, yields surprisingly good results. It is recognized that the true average will be of the minimum soil strengths within a roughly circular band – presumably the area of this band is on average approximated by the area shown in Figure 3.

With an assumed averaging domain, \( A = w \times h \), the geometric average leads to the following definition for \( \bar{c}_u \),

\[
\bar{c}_u = X_g = \exp \left\{ \frac{1}{A} \int_A \ln c_u(x) \, dx \right\}
\]

which, if \( c_u \) is lognormally distributed, is also lognormally distributed. The resulting coefficient

\[
\bar{N}_s = \frac{\bar{c}_u}{\gamma H}
\]

is then also lognormally distributed with mean and variance

\[
\begin{align*}
\mu_{\ln \bar{N}_s} &= \mu_{\ln c_u} - \ln(\gamma H) \\
\sigma^2_{\ln \bar{N}_s} &= \sigma^2_{\ln c_u} = \gamma (w, h) \sigma^2_{\ln c_u}
\end{align*}
\]
The function $\gamma(w, h)$ is the so-called variance function which lies between 0 and 1 and gives the amount that the variance of a local average is reduced from the point value. It is formally defined as the average of correlations between every pair of points in the averaging domain,

$$
\gamma(w, h) = \frac{1}{A^2} \int_A \int_A \rho(\xi - \eta) \, d\xi \, d\eta
$$

Solutions to this integral, albeit sometimes approximate, exist for most common correlation functions. Alternatively, the integral can be calculated accurately using a numerical method such as Gauss quadrature.

The probability of failure, $p_f$, can now be computed by assuming that Taylor’s stability coefficient method holds when using this effective value of cohesion, namely by computing

$$
p_f = P [N_s < N_{\text{crit}}] = \Phi \left( \frac{\ln N_{\text{crit}} - \mu_{\ln N_s}}{\sigma_{\ln N_s}} \right)
$$

where the critical stability coefficient for the 2:1 slope is $N_{\text{crit}} = 0.173$ and for the 1:1 slope is $N_{\text{crit}} = 0.184$. $\Phi$ is the cumulative distribution function for the standard normal. Unfortunately, the geometric average for $\bar{e}_u$ leads to predicted failure probabilities which significantly underestimate the probabilities determined via simulation and changes in the averaging domain size does not particularly improve the prediction. This means that the soil strength as ‘seen’ by the finite element model is even lower, in general, than that predicted by the geometric average. Thus, the geometric average was abandoned as the correct measure for $\bar{e}_u$.

Since the harmonic average yields values which are even lower than the geometric average, the harmonic average over the same domain, $A = w \times h$, is now investigated as representative of $\bar{e}_u$, namely,

$$
\bar{e}_u = X_h = \left[ \frac{1}{A} \int_A \frac{dx}{e_u(x)} \right]^{-1}
$$

Unfortunately, so far as the authors are aware, no relatively simple expressions exist for the moments of $\bar{e}_u$, as defined above, for a spatially correlated random field. The authors are continuing research on this problem but, for the time being, these moments can be obtained by simulation. It may seem questionable to be developing a probabilistic model with the nominal goal of eliminating the necessity of simulation, when that model still requires simulation. However, the moments of the harmonic mean can be arrived at in a small fraction of the time taken to perform the non-linear slope stability simulation.

In order to compute probabilities using the statistics of $\bar{e}_u$, it is necessary to know the distribution of $N_s = \bar{e}_u / (\gamma H)$. For lognormally distributed $e_u$, the distribution of the harmonic average is not simple. However, since $\bar{e}_u$ is strictly non-negative ($e_u \geq 0$), it seems reasonable to suggest that $\bar{e}_u$ is at least approximately lognormal. A histogram of the harmonic averages obtained in the case where $V = 0.5$ and $\theta_{\ln e_u}/H = 0.5$ is shown in Figure 4, along with a fitted lognormal distribution. The $p$-value for the Chi-Square goodness-of-fit test is 0.44, indicating that the lognormal distribution is very reasonable, as also indicated by the plot. Similar results were obtained for other parameter values.

![Figure 4. Histogram of harmonic averages along with fitted lognormal distribution.](image-url)
The procedure to estimate the mean and variance of the harmonic average, $\bar{c}_u$, for each parameter set ($\mu_{N_s}$, $V$, and $\theta_{ln c_u}$) considered in this study involves; a) generating a large number of random cohesion fields, each of dimension $w \times h$, b) computing the harmonic average of each using Eq. (10), and c) estimating the mean and variance of the resulting set of harmonic averages. Using 5000 random field realizations, the resulting estimates for the mean and standard deviation of $\ln \bar{X}_h$ are shown in Figure 5 for random fields with mean 1.0. Since $\bar{c}_u$ is assumed to be (at least approximately) lognormally distributed, having parameters $\mu_{\ln \bar{c}_u}$ and $\sigma_{\ln \bar{c}_u}$, the mean and standard deviation of the logarithm of the harmonic averages are shown in Figure 5 for the two slopes considered.

Figure 5. Mean and standard deviation of log-harmonic averages estimated from 5000 simulations.
Of note in Figure 5 is the fact that there is virtually no difference in the mean and standard deviation for the 2:1 and 1:1 slopes, even though the averaging regions have quite different shapes. Admittedly the two averaging regions have the same area, but this only slow change in harmonic average statistics with averaging dimension has been found also to be true of changing areas. This implies that the accurate determination of the averaging area is not essential to the success of failure probability predictions.

Given the results of Figure 5, the slope failure probability can now be computed as in Eq. (16),

\[ p_f = P[N_s < N_{crit}] = \Phi \left( \frac{\ln N_{crit} - \mu_{\ln N_s}}{\sigma_{\ln N_s}} \right) \]  

(18)

except that now the mean and standard deviation of \( \ln N_s \) are computed using the harmonic mean results of Figure 5, suitably scaled for the actual value of \( \mu_{cu}/\gamma H \) as follows,

\[ \mu_{\ln N_s} = \ln(\mu_{cu}/\gamma H) + \mu_{\ln X_h} = \ln(\mu_{N_s}) + \mu_{\ln X_h} \]  

(19a)

\[ \sigma_{\ln N_s} = \sigma_{\ln X_h} \]  

(19b)

where \( \mu_{\ln X_h} \) and \( \sigma_{\ln X_h} \) are read from Figure 5, given the correlation length and coefficient of variation.

Figure 6 shows the predicted failure probabilities versus the failure probabilities obtained via simulation over all parameter sets considered. The agreement is remarkably good, considering the fact that the averaging domain was rather arbitrarily selected, and there was no a-priori evidence that the slope stability problem should be governed by a harmonic average. The results of Figure 6 indicate that the harmonic average gives a good probabilistic model of slope stability.

Figure 6. Simulated failure probabilities versus failure probabilities predicted using a harmonic average of \( c_u \) over domain \( w \times h \).

There are a few outliers in Figure 6 where the predicted failure probability considerably overestimates that obtained via simulation. For the 2:1 slope, these outliers correspond to the cases where 1) \( \mu_{N_s} = 0.3, V = 1.0 \) and \( \theta_{\ln c_u}/H = 0.1 \) (simulated probability is 0.047 versus predicted probability of 0.86) and 2) \( \mu_{N_s} = 0.3, V = 1.0 \) and \( \theta_{\ln c_u}/H = 0.2 \) (simulated probability is 0.31 versus predicted probability of 0.74). Both cases correspond to the largest factor of safety considered in the study (\( \mu_{N_s} = 0.3 \) gives a factor of safety of 1.77 in the uniform soil case). Also the small correlation lengths yield the smallest values of \( \sigma_{\ln N_s} \) which, in turn, implies that the cumulative distribution function of \( \ln N_s \) increases very rapidly over a small range. Thus, slight errors in the estimate of \( \mu_{\ln N_s} \) makes for large errors in the probability.
For example, the worst case seen in Figure 6(a) has predicted values of

\[ \mu_{\ln N_s} = \ln(\mu_{N_s}) + \mu_{\ln x_h} = \ln(0.3) - 0.66 = -1.864 \]
\[ \sigma_{\ln N_s} = \sigma_{\ln x_h} = 0.10 \]

The predicted failure probability is thus

\[ P[ N_s < 0.173 ] = \Phi \left( \frac{\ln 0.173 + 1.864}{0.10} \right) = \Phi(1.10) \]
\[ = 0.86 \]

As mentioned, a relatively small error in the estimation of \( \mu_{\ln N_s} \) can lead to a large change in probability. For example, if \( \mu_{\ln N_s} \) was \(-1.60\) instead of \(-1.864\), a 14% change, then the predicted failure probability changes significantly to

\[ P[ N_s < 0.173 ] = \Phi \left( \frac{\ln 0.173 + 1.6}{0.10} \right) = \Phi(-1.54) \]
\[ = 0.062 \]

which is about what was obtained via simulation. The conclusion drawn from this example is that small errors in the estimation of \( \mu_{\ln N_s} \) or, equivalently, in \( N_{crit} \) can lead to large errors in the predicted slope failure probability if the standard deviation of \( \ln N_s \) is small. The latter occurs for small correlation lengths, \( \theta_{\ln c_u}/H \). In most cases for small values of \( \theta_{\ln c_u}/H \) the failure probability tends to be either close to zero \((V < 1.0)\) or close to 1.0 \((V > 1.0)\), in which case the predicted and simulated probabilities are in much better agreement. That is, the model shows very good agreement with simulation for all but the case where a large factor of safety is combined with a small correlation length and intermediate coefficient of variation \((V \approx 1.0)\). This means that the selected harmonic average model is not the best predictor in the region where the cumulative distribution is rapidly increasing. However, in these cases, the predicted failure probability is over-estimated, which is at least conservative.

For all other results, especially where the factor of safety is closer to 1.0 \((\mu_{N_s} < 0.3)\), the harmonic average model leads to very good estimates of failure probability with somewhat more scatter seen for the 1:1 slope. The increased scatter for the 1:1 slope is perhaps as expected, since the steeper slope leads to a larger variety of critical failure surfaces. In general for both slopes the predicted failure probability is seen to be conservative at small failure probabilities, slightly overestimating the failure probability.

5. CONCLUSIONS

This study investigates the failure probabilities of two undrained clay slopes, one with gradient 2:1 and the other with gradient 1:1. The basic idea of the paper is that the Taylor stability numbers are still useful if an ‘effective’ soil property can be found to represent the spatially random soil. It was found that a harmonic average of the soil cohesion over a region of dimension \( H^2(\sin \beta \times \frac{1}{\sin \beta}) = H^2 \) yields an effective stability number with an approximately lognormal distribution that quite well predicts the probability of slope failure. The harmonic average was selected because it is dominated by low strength regions appearing in the soil slope, which agrees with how the failure surface will seek out the low strength areas. The dimension of the averaging region was rather arbitrarily selected – the effective stability number mean and variance is only slowly affected by changes in the averaging region dimension – but is believed to reasonably approximate the area of the ‘average’ slope failure band.

An important practical conclusion arising from the fact that soil slopes appear to be well characterized by a harmonic average of soil sample values, rather than by an arithmetic average, as is traditionally done, has to do with how soil samples are treated. In particular, the study suggests that the reliability of an existing slope is best estimated by sampling the soil at a number of locations and then using a harmonic average of the sample values to estimate the soil’s effective cohesion. Most modern geotechnical codes suggest that soil design properties be taken as ‘cautious estimates of the mean’ – the harmonic average, being governed by low strength regions, is considered by the authors to be such a ‘cautious estimate’ for slope stability calculations.
REFERENCES


